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Phil. Trans. R. Soc. B 2013 **368**, 20130166, published 27 May 2013

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Cite this article: Sutton MA *et al.* 2013

Towards a climate-dependent paradigm of ammonia emission and deposition. *Phil Trans R Soc B* 368: 20130166.

<http://dx.doi.org/10.1098/rstb.2013.0166>

One contribution of 15 to a Discussion Meeting Issue 'The global nitrogen cycle in the twenty-first century'.

Subject Areas:

environmental science

Keywords:

ammonia, emission, deposition, atmospheric modelling

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Electronic supplementary material is available at <http://dx.doi.org/10.1098/rstb.2013.0166> or via <http://rstb.royalsocietypublishing.org>.

Towards a climate-dependent paradigm of ammonia emission and deposition

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Existing descriptions of bi-directional ammonia (NH₃) land–atmosphere exchange incorporate temperature and moisture controls, and are beginning to be used in regional chemical transport models. However, such models have typically applied simpler emission factors to upscale the main NH₃ emission terms. While this approach has successfully simulated the main spatial patterns on local to global scales, it fails to address the environment- and climate-dependence of emissions. To handle these issues, we outline the basis for a new modelling paradigm where both NH₃ emissions and deposition are calculated online according to diurnal, seasonal and spatial differences in meteorology. We show how measurements reveal a strong, but complex pattern of climatic dependence, which is increasingly being characterized using ground-based NH₃ monitoring and satellite observations, while advances in process-based modelling are illustrated for agricultural and natural sources, including a global application for seabird colonies. A future architecture for NH₃ emission–deposition modelling is proposed that integrates the spatio-temporal interactions, and provides the necessary

foundation to assess the consequences of climate change. Based on available measurements, a first empirical estimate suggests that 5°C warming would increase emissions by 42 per cent (28–67%). Together with increased anthropogenic activity, global NH₃ emissions may increase from 65 (45–85) Tg N in 2008 to reach 132 (89–179) Tg by 2100.

1. Introduction

Ammonia (NH₃) can be considered as representing the primary form of reactive nitrogen (N_r) input to the environment. In the human endeavour to produce N_r for fertilizers, munitions and other products, NH₃ is the key manufactured compound, produced through the Haber–Bosch process [1]. Synthesis of NH₃ is also the central step in the biological fixation of N₂ to produce organic reduced nitrogen compounds (R-NH₂), such as amino acids and proteins. When it comes to the decomposition of these organic compounds, ammonia and ammonium (NH₄⁺), collectively NH_x, are again among the first compounds produced. These changes lead to a cascade of transformations into different N_r forms, with multiple effects on water, air, soil quality, climate and biodiversity, until N_r is eventually denitrified back to N₂.

Although the behaviour of ammonia has long been of interest at both micro and macro scales [2], recent scientific efforts and policies have given it much less attention than other N_r forms. For example, under revision of the UNECE Gothenburg Protocol in 2012, the controls for NH₃ were the least ambitious of all pollutants considered, with a projected decrease in NH₃ emission for the EU (between 2010 and 2020) of only 2 per cent, compared with reductions of 30 per cent for SO₂ and 29 per cent for NO_x (based on CEIP [3] and UNECE [4]).

In North America, India and China the expected trends are even more challenging. Figure 1 shows the relative changes in atmospheric N_r deposition across the east of North America projected for 2001–2020 [5]. Despite increases in traffic volume, the implementation of technical measures to reduce NO_x emission from vehicles contributes an approximately 40 per cent reduction in oxidized nitrogen (NO_y) deposition. By comparison, the minimal adoption of technical measures to reduce NH₃ emission from agriculture is being offset by increased meat and dairy consumption, requiring more livestock and fertilizers, increasing NH_x deposition in some areas by greater than 40 per cent.

The combination of weak international commitments to mitigate NH₃ and increasing *per capita* consumption represents one of the greatest challenges for future management of the nitrogen cycle [6,7]. The reality is that, rather than needing more N_r to sustain ‘food security’, in developed parts of the world high levels of N_r consumption are being used to sustain ‘food luxury’—the security of our food luxury. Ammonia must be a key part of the societal debate on these issues, where scientific advances in understanding and quantification are essential, especially as NH₃ emission is one of the largest N_r losses.

Most NH₃ emissions result from agricultural production, and are strongly influenced by climatic interactions. In principle, according to solubility and dissociation thermodynamics, NH₃ volatilization potential nearly doubles every 5°C, equivalent to a Q₁₀ (the relative increase over a range of 10°C) of 3–4. At the same time, NH₃ emission is controlled by water availability, which allows NH_x to dissolve, be taken up by organisms and be released through decomposition. Considering these

interactions, NH₃ emission and deposition are expected to be extremely climate-sensitive. For example, will climate warming increase NH₃ emissions and their environmental effects, and to what extent will this hinder NH₃ mitigation efforts?

While substantial advances have been made in process-level understanding of NH₃ land–atmosphere exchange [8–14], these advances have not been fully upscaled at national, continental and global levels. Bi-directional models using the ‘canopy compensation point’ approach [10,15] have only been included to a limited extent in a few chemistry and transport models (CTMs) [5,16–18].

In addition, CTMs are still largely based on precalculated emission inventories. Under this approach, activity statistics are combined with emission factors to estimate annual emissions, which are mapped typically with relatively simple temporal disaggregation. The resulting fixed emission estimates are attractive to policy users in relation to reporting national emissions commitments. However, the approach fails to recognize that a warm-dry year would tend to give larger NH₃ emissions than a cold-wet year. At the same time, it does not address the short-term interactions relevant for risk assessment of NH_x impacts [5,19,20].

To address these issues, this paper examines the relationships between climatic drivers and ammonia exchange processes. We first consider the magnitude of global NH₃ emissions. Following consideration of the process relationships controlling NH₃ exchange, we show how studies of a natural NH₃ source (seabird colonies) can be used to demonstrate the climatic dependence of emissions and verify a global model. Finally, we outline a new architecture that sets the challenge for a new paradigm for regional modelling of atmospheric NH_x as the basis for incorporating the effects of climate differences and climate change.

2. Ammonia emission inventories

The main reasons for constructing NH₃ emissions inventories have been to meet national-scale policy requirements and provide input to CTMs. Among the best studied national NH₃ inventories are those of the Netherlands [21], Denmark [22], the UK [23,24], Europe [25] and the US [5].

Although there is frequent debate on the absolute magnitude of national emissions and their consistency with atmospheric measurements [26], such inventories have allowed high-resolution CTMs to show a close spatial correlation with annual atmospheric NH₃ and NH₄⁺ concentrations. In Europe, the inventories have focused especially on livestock housing and grazing, storage and spreading of manures, and from mineral fertilizers [27]. Less attention has been given to non-agricultural emissions including sewage, vehicles, pets, fish ponds, wild animals and combustion, which can contribute 15 per cent to national totals [28,29].

By comparison with the best national estimates, global NH₃ emission inventories are much less certain. This reflects the wider diversity of sources and fewer underpinning data, combined with a paucity of activity statistics (e.g. animal numbers, bodyweights, diets, etc.). The contrast is illustrated between Denmark, where 1 km resolution data on livestock numbers account for species sub-classes and abatement techniques [30], and other parts of the world, where such statistics often do not even exist.

Recent global estimates of annual NH₃ emission are summarized in table 1. Dentener & Crutzen [16] were the first to

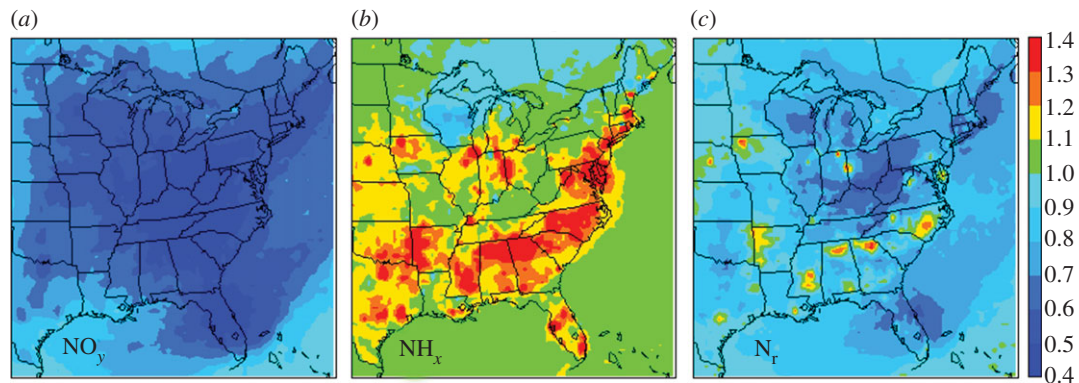


Figure 1. Simulated changes in N deposition in eastern USA, showing the ratios for 2020/2001 (adapted from Pinder *et al.* [5]). (a) Oxidized N deposition, (b) reduced N deposition and (c) total N deposition.

Table 1. Comparison of global ammonia emission estimates (Tg N yr^{-1}).

	Dentener & Crutzen [16]	Bouwman <i>et al.</i> [31]	Van Aardenne <i>et al.</i> [32]	Beusen <i>et al.</i> [33]	PBL/JRC [34] EDGAR 4.2	PBL/JRC [34] EDGAR 4.2	current best estimates (total all sources)	
year	1990	1990	2000	2000	2000	2008	2000	2008
spatial resolution	$10^\circ \times 10^\circ$	$1^\circ \times 1^\circ$	$1^\circ \times 1^\circ$	$0.1^\circ \times 0.1^\circ$	$1^\circ \times 1^\circ$	$1^\circ \times 1^\circ$		
excreta from domestic animals	22.0 ^a	21.6 ^a	21.1 ^a	21.0 ^a	8.0 ^b	8.7 ^b	8.0 ^b	8.7 ^b
use of synthetic N_r fertilizers	6.4	9.0	12.6	11.0				
agricultural soils and crops	—	3.6 ^c	—	—	21.6 ^a	24.7 ^a	25.2 ^{a,c}	28.3 ^{a,c}
biomass burning ^d	2.0	5.9	5.4	—	4.4	5.5	4.4	5.5
industrial and fossil fuel burning ^e	—	0.3	0.3	—	1.3	1.6	1.3	1.6
human population and pets ^f	—	2.6	—	—	—	—	3.0	3.3
waste composting & processing ^f	—	—	4.0	—	0.01	0.02	4.0	4.4
soils under natural vegetation	5.1	2.4	—	—	—	—	2.4	2.4
excreta from wild animals ^g	2.5	0.1	—	—	—	—	2.5	2.5
oceans (and volcanoes)	7.0	8.2	—	—	—	—	8.6 ^h	8.6 ^h
total	45.0	53.6	43.0	—	35.2	40.6	59.3	65.4
total from livestock and crops	28.4	34.2	33.7	32.0	29.6	33.4	33.2	37.0

^aIncludes emissions from grazing and land application of animal manure.

^bExcludes emissions from land application of animal manure.

^cIncludes estimated direct crop emissions from foliage and leaf litter.

^dIncluding savannah, agricultural waste, forest, grassland and peatland burning/fires.

^eNot including potentially high emissions from low-efficiency domestic coal burning [2].

^fRescaled by global population increase.

^gThe estimate of Bouwman *et al.* [31] is considered low given NH_3 emissions from seabird colonies alone of 0.3 Tg N yr^{-1} [35].

^hIncludes an upper estimate of 0.4 Tg N yr^{-1} as NH_3 from volcanoes based on an emission ratio of 15% NH_3 : SO_2 [2] and volcanic SO_2 emissions of 6.7 Tg S yr^{-1} [36].

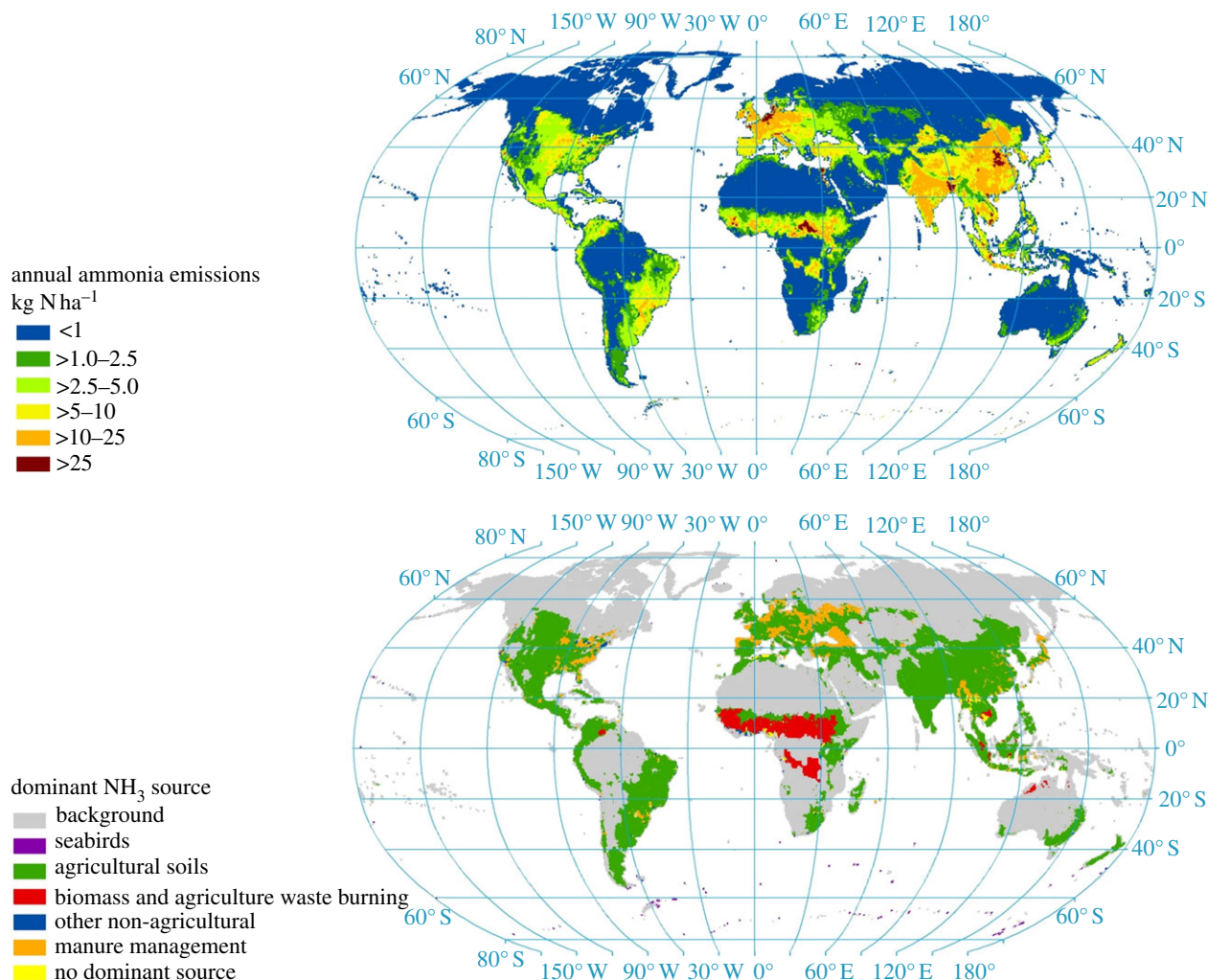


Figure 2. Spatial variability in global ammonia emissions based on JRC/PBL [34] (livestock, fertilizers, biomass burning, fuel consumption) and Riddick *et al.* [35] (seabirds). Emissions from oceans, humans, pets, natural soils and other wild animals (table 1) are not mapped. High-resolution maps for the UK are given in the electronic supplementary material, figure S1.

derive a global $10^\circ \times 10^\circ$ NH_3 emission inventory for input to global CTMs. Bouwman *et al.* [31] made a global NH_3 inventory for the main sources at $1^\circ \times 1^\circ$ for 1990, while Beusen *et al.* [33] extended this for livestock and fertilizers.

One of the first points to note in the global comparison is that the source nomenclature is not well harmonized. Current standardization of inventory reporting by EDGAR (Emission Database for Global Atmospheric Research [34]) and the UNFCCC (United Nations Framework Convention on Climate Change) focuses strongly on combustion sources and is less suited for sector analysis of agricultural emissions. It is, therefore, not easy to distinguish the main livestock sectors in the most recent inventories. According to Dentener & Crutzen [16], of 22 Tg $\text{NH}_3\text{-N yr}^{-1}$ emitted from livestock, 65 per cent was from cattle and buffalo, with 13 per cent, 11 per cent, 6 per cent and 5 per cent, from pigs, sheep/goats, poultry and horses/mules/asses, respectively.

The degree of agreement shown in table 1 (35–54 Tg N yr^{-1}) results partly from dependence on common datasets (e.g. Food and Agriculture Organization) and partly because of including different emission terms in each inventory. If all sources listed among the inventories are combined, this gives a total of 59 and 65 Tg N yr^{-1} for 2000 and 2008, respectively. These values are based on the recent estimates of EDGAR, combined with approximately 8 Tg yr^{-1} from oceans and

approximately 12 Tg yr^{-1} from humans, waste, pets, wild animals and natural soils.

These estimates should be considered uncertain by at least $\pm 30\%$ (based on propagation of likely ranges for input data, [33]), indicating an emission range of 46–85 Tg N for 2008, although a formal uncertainty analysis on the full inventory has never been conducted. Apart from the uncertainties related to emission factors and climatic dependence, inaccurate activity data may introduce regional bias. For example, comparison of NH_3 satellite observations (see §4) with a global CTM showed substantial underestimation by the CTM in central Asia [37], suggesting an under-reporting of animal numbers and fertilizer use in these countries.

Figure 2 shows that the regions of the world with highest emissions are mostly associated with livestock and crops. Because the available sector categorization does not distinguish arable and livestock sectors, the orange-shaded areas represent locations with a very strong livestock dominance. Biomass burning is the main NH_3 source across much of central Africa, where estimated NH_3 emissions reach levels similar to peak agricultural values of India and China. Inclusion of the recent estimates of Riddick *et al.* [35] shows how seabird colonies are a significant NH_3 source for many subpolar locations. These global maps hide substantial local variability, as illustrated for the UK in the electronic supplementary material, figure S1.

It must be emphasized, however, that these global estimates only take climate factors into account in a limited way. For emissions from fertilizer and manure application, climate has been partly considered by grouping datasets into major temperature regions [38], whereas Riddick *et al.* [35] applied a simple temperature function. However, the published global inventories do not model NH_3 at a process level in relation to changing meteorological conditions. In addition, bi-directional NH_3 fluxes from crops, sparsely grazed land and natural vegetation provide a particular challenge, because both the magnitude and direction of the flux varies according to ecosystem, management and environmental variables.

3. Concepts for modelling ammonia land–atmosphere exchange

Current conceptual frameworks on NH_3 exchange show how fluxes respond to short-term variation in environmental conditions, and hence to long-term climate differences. This can be illustrated by the case of bi-directional exchange between plant, soil and atmosphere.

Ammonia fluxes are often considered as representing a potential difference between two gas-phase concentrations constrained by a set of resistances. At its simplest, the concentration at the surface $\chi(z_{o'})$, where $z_{o'}$ is the notional height of NH_3 exchange, is contrasted with the concentration $\chi(z)$ at a reference height z above the canopy, with the total flux (F_t):

$$F_t = \frac{[\chi(z_{o'}) - \chi(z)]}{[R_a(z) + R_b]}, \quad (3.1)$$

where $R_a(z)$ and R_b are the turbulent atmospheric and quasi-laminar boundary layer resistances, respectively [10,15]. A well-known variant of this approach, applicable only for deposition, assumes that the concentration at the absorbing surface is zero, so that any limitation to uptake can be assigned to a canopy resistance (R_c):

$$F_t = \frac{[0 - \chi(z)]}{[R_a(z) + R_b + R_c]}, \quad (3.2)$$

where an associated term, the deposition velocity, is defined as $V_d(z) = (R_a(z) + R_b + R_c)^{-1} = -F_t/\chi(z)$. It is possible to interpret NH_3 flux measurements according to either view. This is illustrated in the electronic supplementary material, figure S2, which summarizes results from a year of continuous hourly NH_3 flux measurements over an upland moorland in Scotland [39]. Applying equation (3.1) to the flux measurements demonstrates the relationship between $\chi(z_{o'})$ and canopy temperature, while applying equation (3.2) to calculate R_c for the same dataset is necessarily restricted to periods where deposition was recorded. These two approaches represent different views of the factors driving and constraining the net flux.

The value of $\chi(z_{o'})$ at the surface is proportional to a ratio termed $\Gamma = [\text{NH}_4^+]/[\text{H}^+]$, where according to the thermodynamics:

$$\chi = \frac{161\,500}{\text{Texp}^{(-10\,380/T)} [\text{NH}_4^+]/[\text{H}^+]}, \quad (3.3)$$

with T in Kelvin [15]. The existence of bi-directional fluxes illustrated in the electronic supplementary material, figure S2, shows that calculating $\chi(z_{o'})$ provides the more general solution, whereas its increase according to thermodynamics (fitted line, $Q_{10} = 3\text{--}4$) suggests that it reflects a process

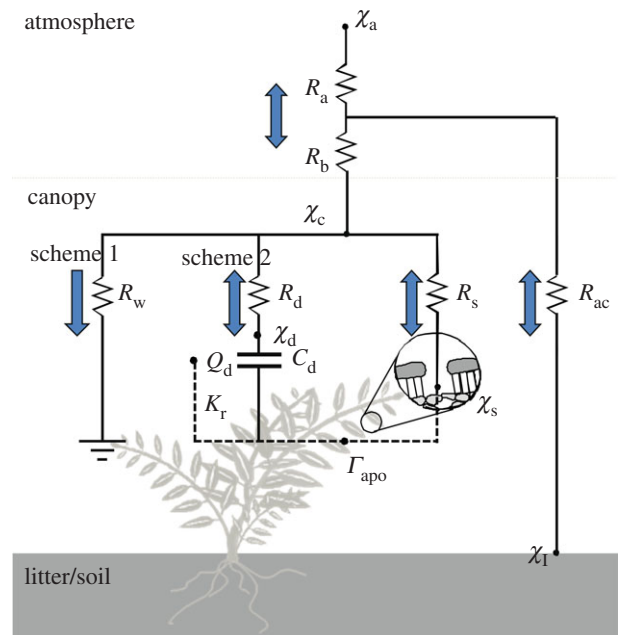


Figure 3. Resistance analogue of NH_3 exchange including cuticular, stomatal and ground pathways. Two schemes for cuticular exchange are illustrated: scheme 1, steady-state uptake according to a varying cuticular resistance (R_w); scheme 2, dynamic exchange with a pool of NH_4^+ treated with varying capacitance (C_d) and charge (Q_d). Other resistances are for turbulent atmospheric transfer (R_a), the quasi-laminar boundary layer (R_b), within-canopy transfer (R_{ac}), cuticular adsorption/desorption (R_d) and stomatal exchange (R_s). Also shown are the air concentration (χ_a), cuticular concentration (χ_d), stomatal compensation point (χ_s), litter/soil surface concentration (χ_l) and the canopy compensation point (χ_c). Exchange between aqueous NH_4^+ pools is shown with dashed lines, including K_r , the exchange rate between leaf surface and apoplast.

reality. An exception is seen in frozen conditions, where R_c may be better suited to describe slow rates of deposition, as seen also for other gases [40]. However, considering the full year of measurements, the clear relationship with $\chi(z_{o'})$ in electronic supplementary material, figure S2, illustrates the weakness of sole reliance on the R_c and V_d approach typically applied in CTMs.

The approach described above outlines the most simple situation. In reality, each of surface concentrations, resistances and even capacitances can be used to simulate NH_3 exchange, whereas both advection and gas–particle interactions can also affect fluxes [11,41,42]. A framework to consider the key issues at the plot scale is shown in figure 3, largely based on Sutton *et al.* [10,43], Flechard *et al.* [44] and Nemitz *et al.* [15]. In this development of the resistance analogy, the central term is the ‘canopy compensation point’ (χ_c), which is identical to $\chi(z_{o'})$ when $F_t = 0$. This is contrasted with the ‘stomatal compensation point’ (χ_s), which is the NH_3 gas concentration at equilibrium with $[\text{NH}_4^+]/[\text{H}^+]$ in the apoplast, Γ_{apo} . Available data suggest only modest diurnal variation in Γ_{apo} [11]. The main challenge, therefore, is to estimate the larger differences in Γ_{apo} owing to management, plant species and seasonality [45–47]. This can be investigated by including NH_3 cycling in models of ecosystem dynamics and agricultural management [11,18,48–50].

The most widely used approach to simulate NH_3 exchange with the cuticle is to assume that deposition is constrained by a cuticular resistance (R_w) [10,15]. General parametrizations of this response to humidity and to NH_3 :

acid-gas ratios have been developed ([15,46,51]; electronic supplementary material, figure S3; figure 3, scheme 1). These approaches have the advantage of relative simplicity, but only represent a steady-state approximation to a dynamic reality, where both adsorption (favouring net deposition) and desorption (favouring emission) occur in practice.

This dynamic view can be addressed by scheme 2 of figure 3. In the simplest description, a time-constant can be set for charging and discharging the leaf-surface water/cuticular pool of NH_x (e.g. $R_d = 5000/C_d$, s m^{-1}), combined with a fixed leaf-surface pH [43]. A more sophisticated approach solves the ion balance of the leaf-surface water, calculating cuticular pH according to the concentrations, fluxes and precipitation inputs of all relevant compounds [44,52]. In an extended application to measurements over forest, Neirynck & Ceulemans [53] tested the simpler application of scheme 2, finding it to simulate diurnal to seasonal measured fluxes much more closely than scheme 1. One of the main uncertainties in applying scheme 2 is the exchange of aqueous NH_4^+ and other ions between leaf surface and apoplast.

The last component of figure 3 describes NH_3 exchange with the ground surface. Although flux measurements have often shown significant emission from soils and especially with leaf litter [8,11,15], this term remains the most uncertain. In particular, the extent to which soil pH influences pH at atmospheric exchange surfaces such as leaf litter and in the vicinity of applied fertilizer and manure remains poorly quantified, while the liberation of NH_3 from organic decomposition directly influences local substrate pH. Further measurements are needed to develop simple parametrizations of Γ_l for litter and to inform the development of ecosystem models simulating Γ_l in relation to litter quality, water availability, mineralization, immobilization and nitrification rates.

Similar interactions apply to other NH_3 volatilization sources. For example, the VOLT'AIR model provides a process simulation of NH_3 emissions from the land application of liquid manures [54,55], where manure placement method and calculated soil infiltration rates inform the calculation of $\chi(z_o)$. Empirical approaches have also been used to parametrize NH_3 emissions directly from manure application, using regression with experimental studies [56]. Such empirical relationships have also been applied to estimate NH_3 emissions from animal houses, manure stores and manure spreading (see the electronic supplementary material, section S3). It remains a future challenge to develop process models for these sources based on the principles of equation (3.1).

4. Quantifying environmental relationships with ammonia fluxes and concentrations

From the preceding examples it can be seen that temperature and moisture play a key role in determining the concentration of NH_3 in equilibrium with surface pools and hence in defining net NH_3 fluxes on diurnal to annual scales. However, the interaction between these and other factors (e.g. stomatal opening, growth dilution of NH_x pools, soil infiltration and decomposition rates) means that the temperature-dependence of NH_3 emission may not always follow the thermodynamic response.

These interactions are illustrated in the electronic supplementary material, figure S3b, which shows the values of χ_c for periods when the net flux was zero, from NH_3 flux

measurements over dry heathland [57]. Under very dry conditions (relative humidity, $h < 50\%$), cuticular fluxes appear to have been small, so that $\chi_c \sim \chi_s$, with a clear temperature dependence in the range of Q_{10} 2–4. By contrast, at h in the range 50–70%, there was less relationship to temperature, pointing to a significant role of cuticular adsorption/desorption processes.

The way in which plant growth interactions may alter the temperature response of NH_3 fluxes can be illustrated by the process model PaSim. Based on its application to measured fluxes over Scottish grassland [48], the model was used to consider scenarios of altered annual air temperature, keeping all other factors the same as the original simulation. The effect of cutting and fertilization on the net NH_3 flux and illustrative N_r pools are shown in figure 4. PaSim included the standard thermodynamic dependence of both χ_s and χ_{soil} (equation (3.3)) to simulated values of Γ_s and Γ_{soil} (combined with scheme 1 using $R_w = 30e^{(100-h)/7} \text{ s m}^{-1}$). It is, therefore, notable that the net flux showed only a modest temperature-dependence, with the net flux for the month increasing with $Q_{10} = 1.5$. Only immediately after fertilizer application did the flux increase with $Q_{10} = 3.2$. This can be explained by the warmer temperatures leading to more rapid grass growth, decreasing leaf substrate N_r and modelled Γ_s (figure 4). Although further measurements have also shown a role of leaf litter processes not currently treated in PaSim [11], the simulation demonstrates how growth-related factors can offset the simple thermodynamic NH_x response.

A similar message emerges for NH_3 emissions from land-spreading of manures using the ALFAM multiple regression model [56], which includes a weak temperature response ($Q_{10} = 1.25$). This broadly agrees with a simple empirical approach for pig slurry for the UK ammonia inventory [27], though, for cattle slurry, the distinction between summer and other months in the inventory equates to $Q_{10} = 2.5$. In this case, the key interaction appears to be between volatilization potential and slurry infiltration rate, which can be limited in both waterlogged and hard-dry soils.

Atmospheric NH_3 monitoring can also inform the simulation of seasonal dynamics. In the case of the UK and Danish ammonia networks, areas dominated by cattle and pig show peak emissions in spring, which are reproduced by models accounting for the timing of manure spreading [30,58]. However, the UK also includes substantial background areas (see the electronic supplementary material, figure S1) with a pronounced summer maximum and winter minimum of 0.43 and 0.04 $\mu\text{g NH}_3 \text{ m}^{-3}$, respectively, while sheep-dominated upland areas show a similar annual cycle (0.95 and 0.17 $\mu\text{g m}^{-3}$, respectively). These seasonal patterns are not reproduced in CTMs as they do not adequately treat the climatic dependence of grazing emissions [58]. As the grazing animals that dominate emissions in these areas are outdoors all year, a 12°C difference between the mean temperature of warmest and coolest month equates to a Q_{10} of 9.0 and 4.7 for background and sheep sites, respectively. These large values suggest that other factors enhance the temperature-dependence of NH_3 concentrations, with more rapid scavenging in winter.

Such seasonal differences can also be seen from globally monitored satellite columns of NH_3 at $12 \times 12 \text{ km}^2$ resolution at nadir, through processing of retrievals from the infrared atmospheric sounding interferometer on the MetOp platform. This approach is based on the absorption spectra of NH_3 in the infrared and depends on a strong thermal contrast between

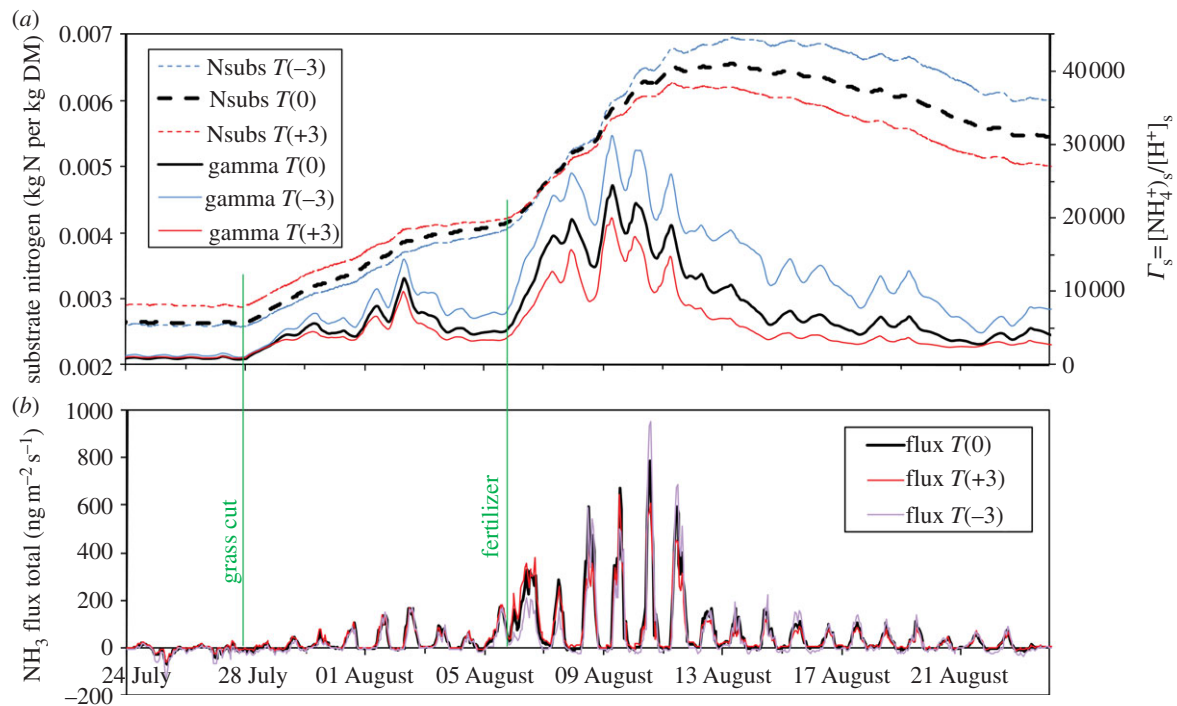


Figure 4. Effect of temperature scenarios (annual change of +3°C and −3°C) on (a) simulated nitrogen pools (foliar substrate N, and $I_{\text{NH}_3}^{\text{f}}$) and (b) net NH₃ fluxes. Simulations conducted using the PaSim model for managed grassland in Scotland following cutting and fertilization with ammonium nitrate.

the ground and atmosphere, measuring NH₃ columns that are dominated by high concentrations in the lowest 1–2 km [37]. Retrievals are made twice a day, allowing extensive comparison with environmental and seasonal NH₃ dynamics.

An illustration of the satellite retrieval is shown in figure 5, which compares the mean NH₃ column over Europe with the seasonally varying NH₃ column at three sites where ground-based monitoring of NH₃ concentrations is available. The map distinguishes areas of high agricultural NH₃ emissions in Brittany, E England, the Netherlands and NW Germany, Po Valley and Nile Delta, while showing high values across Belarus and SW Russia related to forest fires during 2010. The magnitude of the NH₃ columns are also a function of spatial differences in atmospheric mixing that might explain why smaller values are seen in the west compared with the east of the UK. For Stoke Ferry, where NH₃ emissions are dominated by pig and poultry (see the electronic supplementary material, figure S1), both the ground-based and satellite data show spring peak NH₃ values, associated with land-spreading of manure. At Vredepeel, an area of intense pig and cattle farming in the Netherlands, there is less seasonality in the NH₃ data, indicating a stronger contribution of controlled environment livestock housing. Lastly, at K-Puszt, a Hungarian site more distant from local sources, NH₃ levels are highest in summer and lowest in winter, reflecting the integration of different environmentally dependent sources.

The satellite approach requires a strong thermal contrast, limiting its capability in winter and cloudy conditions. However, it allows the examination of spatial patterns and temporal trends with a global coverage that could never be achieved by ground-based air sampling. It thus provides an unprecedented opportunity to improve our understanding of the sources, management and climate controls on NH₃, as further illustrated by seasonal NH₃ patterns in different parts of the world. In the case of the Po Valley, Nile Delta, California and Pakistan, there is a strong seasonal cycle in NH₃, with values of Q_{10} of the column totals mostly in the

range 2–3. However, not all locations show such a temperature-dependence, especially where management differences drive seasonality in NH₃ emissions as seen in livestock dominated areas of Belgium and China (see the electronic supplementary material, figure S4). In order to derive the maximum value from the satellite data, these, therefore, need to be interpreted using detailed atmospheric models, as a basis to disentangle the different driving factors.

5. Seabirds as a model system to assess climate-dependence of global ammonia emissions

The preceding examples highlight the many factors controlling NH₃ emissions, including management effects. In the case of monitoring NH₃ concentrations and atmospheric columns, an even larger number of meteorological factors affect observed values. For these reasons, there is a strong case to use model ecosystems to assess the climate-dependence of NH₃ exchange. At present, this can uniquely be demonstrated by the case of NH₃ emissions from seabird colonies, building on recent measurements and modelling [35,59].

Seabird colonies provide several advantages as a 'model system' to investigate the climate-dependence of NH₃ emissions: the birds follow a well-established annual breeding cycle little affected by human management; rates of N_r excretion can be directly related to dietary energetics for well-characterized populations; and they typically form locally strong NH₃ sources in areas of low NH₃ background. Riddick *et al.* [35] estimated global NH₃ emission from seabird colonies at 0.3 (0.1–0.4) Tg yr⁻¹. Although this is a small fraction of total emissions, it includes major point/island sources greater than 15 Gg NH₃ yr⁻¹, with sites distributed globally across a wide range of climates.

Colony-scale NH₃ flux measurements from seabird colonies were first reported by Blackall *et al.* [59] for Scottish islands, and these have been extended for contrasting climates

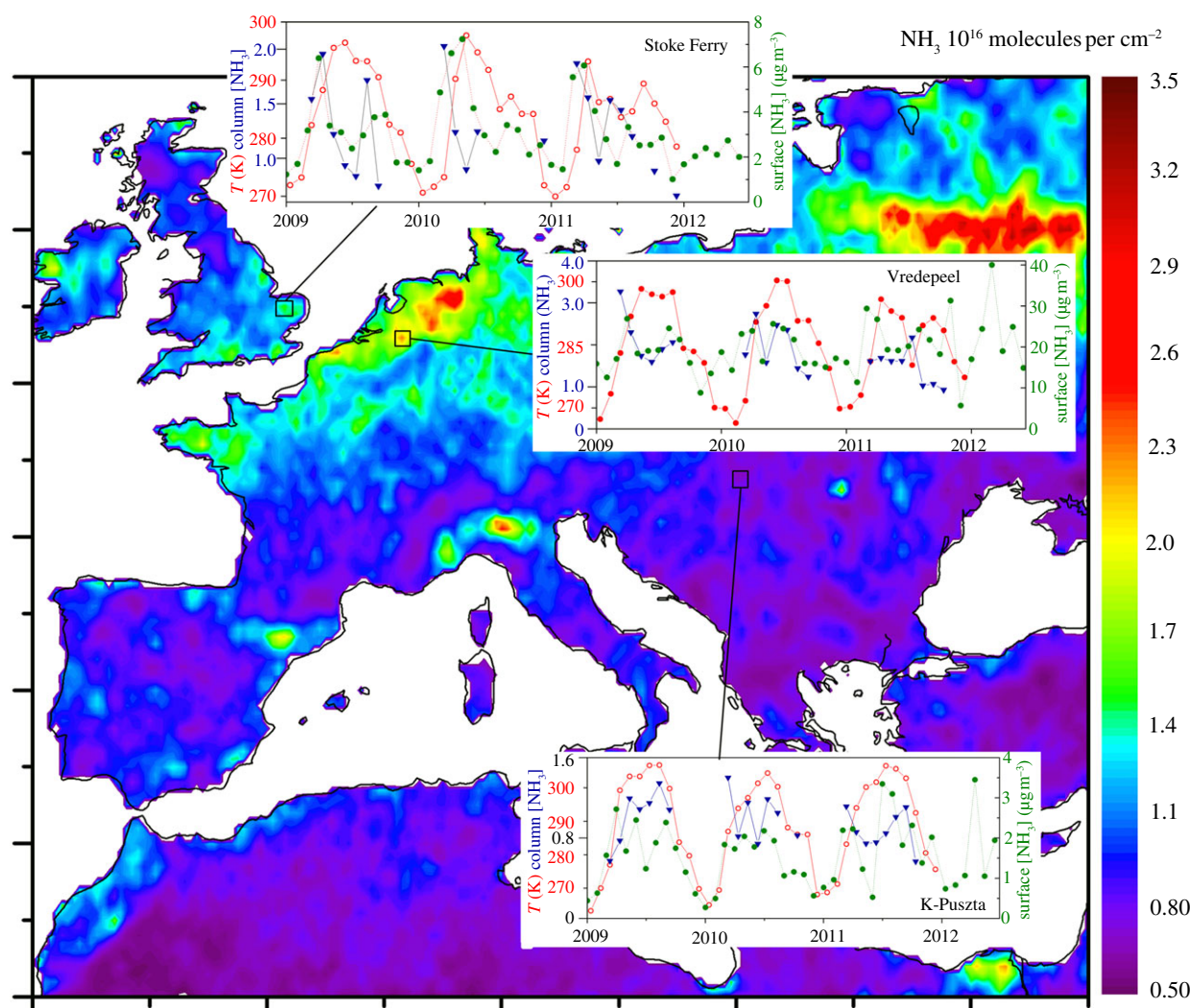


Figure 5. Satellite estimates of the NH_3 column (10^{16} molecules cm^{-2}) and ground temperature, showing the mean for 2009, 2010 and 2011 (from the infrared atmospheric sounding interferometer on the MetOp platform), as compared with ground-based measurements of atmospheric NH_3 concentrations at three selected sites.

as illustrated in figure 6. In this graph, measured NH_3 emissions have been normalized by calculated N_r excretion rates to show the percentage of N_r that is volatilized (P_v). The measurements show a clear temperature-dependence across the globe, with $Q_{10} \sim 3$. For comparison, the dotted line is the estimate used by Blackall *et al.* [59] for global upscaling, whereas the solid line is the initial temperature-adjusted upscaling of Riddick *et al.* [35], following equation (3.3) (their scenario 2).

The importance of these measurements is emphasized by their use to verify a process-based model of NH_3 emissions, the GUANO model (see the electronic supplementary material, figure S5). The model is driven by excretal inputs according to bird diet, energetics and numbers combined with a water-balance to estimate liquid-phase N_r concentrations and run-off. Hydrolysis of uric acid to ammoniacal nitrogen is moisture- and temperature-dependent. By combining the modelled value of $[\text{NH}_4^+]$ with a guano pH of 8.5 and ground surface temperature, equation (3.3) allows estimation of $\chi(z_0)$. This is then applied in equation (3.1) to calculate hourly NH_3 emission.

Application of the GUANO model shows close agreement with measurements, the hourly NH_3 fluxes responding to fluctuations in surface temperature, precipitation events and wind speed. The overall measured temperature-dependence is also reproduced by the GUANO model (figure 6),

including a difference between the two warmest sites, Michelmas Cay and Ascension Island. This is explained by the latter being very dry, limiting rates of uric acid hydrolysis, and hence both measured and modelled NH_3 emission.

Based on the verification of the GUANO model with field measurements, the global seabird and excretion datasets [35] have been applied in the model for hourly simulation of 9000 colonies for 2010–2011 (figure 7). Ground temperature turns out to be the primary driver globally, with P_v ranging from 20 to 72 per cent for sites with annual mean temperature of 30°C , whereas for sites with a mean temperature of 0°C , P_v was 0–18%. Variation between sites of similar temperatures is mainly attributable to differences in water availability, wind speed and nesting habitat (e.g. bare rock versus burrow breeders).

6. Climate-dependent assessment of ammonia emissions, transport and deposition

The examples presented for terrestrial systems including grassland, shrubland, forest and seabird colonies demonstrate the clear climatic dependence of NH_3 exchange processes. Agricultural systems are more complex, and include interactions with management (including alteration of management timing and

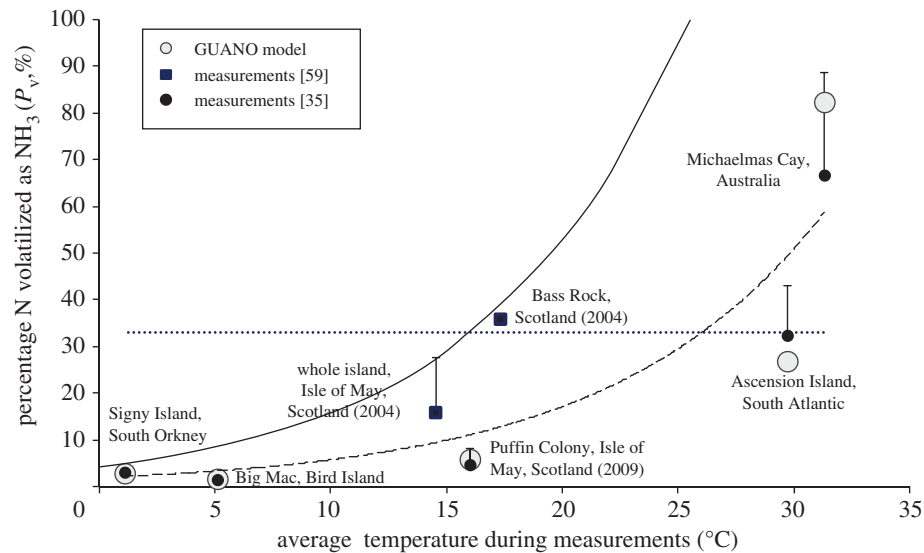


Figure 6. Measured percentage of excreted N_r that is volatilized as NH_3 (P_v) as a function of mean temperature during field campaigns (dashed line: $P_v(\%) = 1.9354e^{0.109T}$; $R^2 = 0.75$), as compared with estimates from the GUANO model for a global selection of seabird colonies. The dotted line shows the value used in a first bioenergetics (BE) model of Blackall *et al.* [59], while the solid line was applied in a temperature-adjusted bioenergetics (TABE) model, by Riddick *et al.* [35] using equation (3.3). The bars on the measured points apply to colonies including burrow-nesting birds and indicate the estimated P_v if the colony were entirely populated by bare-rock breeders.

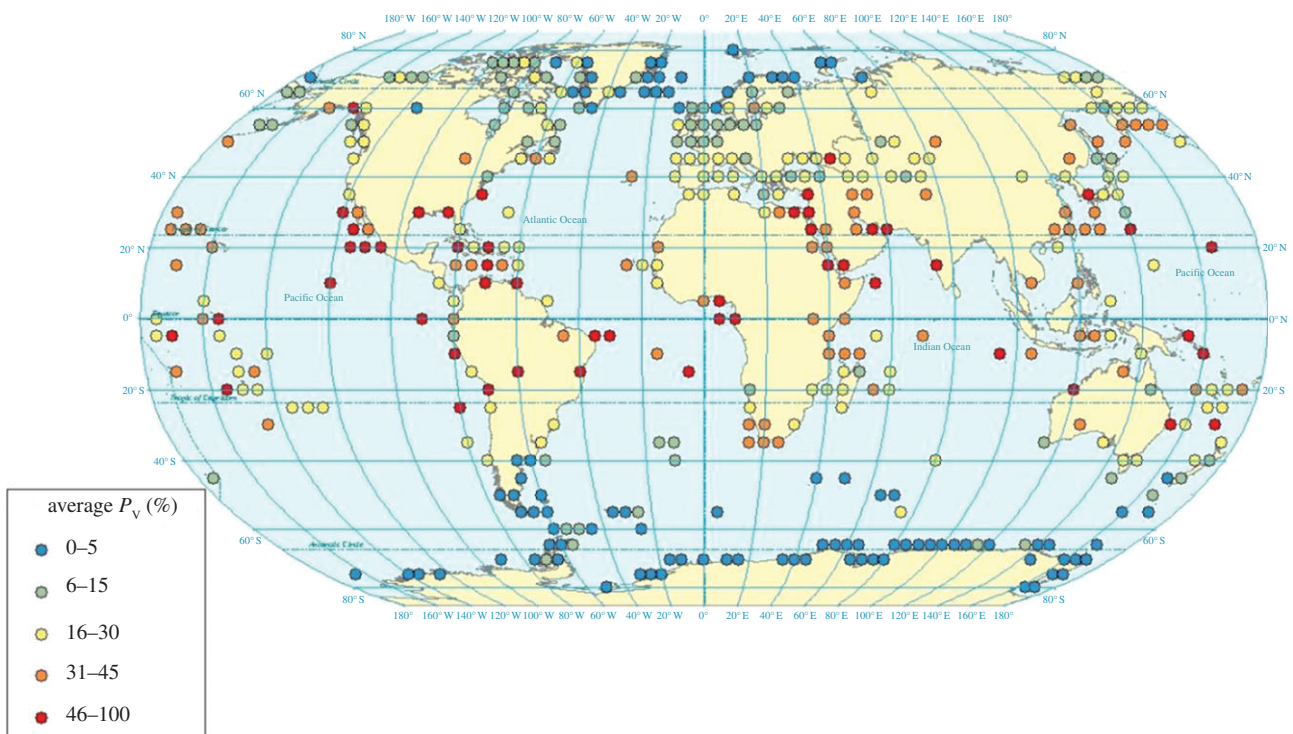
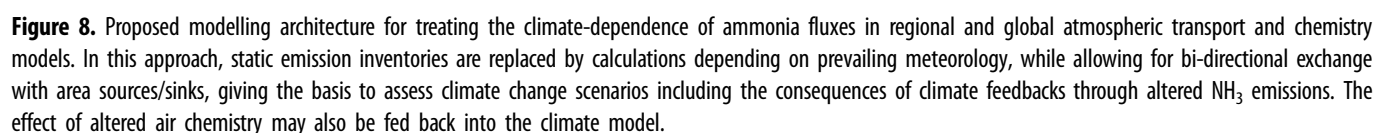


Figure 7. Global application of the GUANO model illustrating the average percentage of excreted N that is volatilized as NH_3 . Excretion calculated based on colony seabird energetics [35], combined with hourly meteorological estimates through 2010–2011.

systems), N_r type, animal housing and manure application method. In principle, however, many of the climatic interactions apply, and can be addressed using process-based models. The same is true for ocean–atmosphere NH_3 exchange, which is bi-directional according to equation (3.1), with $\chi(z_o)$ depending on variations in sea surface temperature, $[NH_4^+]$ concentration, water pH and local NH_3 air concentrations. For example, future ocean acidification would tend to decrease sea-surface NH_3 emission. Of these factors, Johnson *et al.* [60] found temperature to be of overriding importance in determining ocean NH_3 emissions, through its control of $\chi(z_o)$.

With this background, we return to the question of regional and global modelling of NH_3 emission, dispersion and deposition in CTMs. Section 2 showed that there are several limitations in current NH_3 emission inventories, such as information on activity data (numbers and location of animals, fertilizers, fires, etc.), average emission rates and data structure (distinction of source sectors). On a global scale, however, and given the target to assess climate change effects, by far the main limitation is that current architecture uses previously calculated emissions as input to CTMs. In reality, the same meteorology incorporated within a CTM to describe chemical



An outline of the proposed modelling architecture is given in figure 8, with the key new elements highlighted in green. Instead of activity data and experimentally derived relationships being used directly to provide an ‘emissions inventory’, with subsequent (uni-directional) dry deposition, emissions are treated in two submodels: (i) uni-directional emissions from point sources such as manure storage facilities and animal housing (where $\chi(z_o') \gg \chi(z)$) and (ii) bi-directional fluxes from area sources (where $\chi(z_o')$ is less than or greater than $\chi(z)$), which includes emissions or dry deposition according to prevailing conditions. The same meteorological data are thus used to

Although not all these linkages have yet been made, significant progress in the temporal distribution of NH_3 emissions according to agricultural activities has already been achieved, which can provide key input to the future developments [30,61,62]. For example, the US EPA Community Multiscale Air-Quality model includes coupling to an agro-ecosystem model to provide dynamic and meteorological-dependent

emissions from fertilizer application, using a two-layer bidirectional resistance model based on Nemitz *et al.* [15] that includes the effects of soil nitrification processes [18]. Similarly, Hamaoui-Lagel *et al.* [55] incorporated the VOLT' AIR model to simulate NH₃ emissions from fertilizer application in a regional-scale atmospheric model.

The consequences of such temporal interactions can be illustrated by the comparison of measured NH₃ concentration and simulations of a Danish model [22] at a long-term monitoring site (Tange, electronic supplementary material, figure S6). In this case, the model has been used to provide the temporal disaggregation of previously calculated annual emissions. The challenge for the next stage must be to incorporate the environmental drivers in process models for all major sources to quantify the dynamics on hourly, diurnal, seasonal and annual scales, and as a foundation to estimate the effects of long-term climate change.

7. Conclusions

This paper has shown how ammonia emissions and deposition are fundamentally dependent on environmental conditions. While temperature has been found to be the primary environmental driver, other key factors include interactions with canopy and soil wetness and with management practices for agricultural sources. For several systems, such as emission from manure spreading, fertilizers, seabird colonies and bidirectional exchange with vegetation, process models are already available that describe the key relationships.

A new paradigm for atmospheric modelling of NH₃ is proposed, where process models are incorporated with the relevant statistical data to simulate NH₃ emissions as part of atmospheric models. Seabird colonies have been used here to demonstrate the global application of such a process model, verified by measurements under different climates, where the fraction of available N_r volatilized as NH₃ can increase by a factor of more than 20 between subpolar and tropical

conditions. Although a few CTMs have incorporated bidirectional exchange, work is required to parametrize models for different ecosystem types and climates, and to assess the consequences of different levels of model complexity, including the coupling with ecosystem and agronomic models.

The proposed developments provide the necessary foundation to assess how climate will affect NH₃ emissions, dispersion and deposition. The practical implications are that inventory activities should focus increasingly on supplying the statistical activity data needed to drive the models (rather than only publishing static NH₃ emission estimates) and that national NH₃ emissions for any year can only be calculated with confidence once the meteorological data are available.

Based on the available measurements and models, it is possible to indicate empirically the scale of the climate risk for NH₃. Marine NH₃ emissions are expected to follow the thermodynamic response directly (equation (3.3)), whereas a reduced Q_{10} of 2 (1.5–3) may be applied for terrestrial volatilization sources. (For procedures, see the electronic supplementary material, section 8, figures S7, S8 and equations for use in scenario models.) Applying these responses to the 2008 global estimates of 65 (46–85) Tg N yr⁻¹ for a 5°C global temperature increase to 2100 would increase NH₃ emissions by approximately 42 per cent (28–67%) to 93 (64–125) Tg. If this is combined with a further 56 per cent (44–67%) increase in anthropogenic source activities [63,64], total NH₃ emissions would reach 132 (89–179) Tg by 2100. Considering these major anticipated increases, the limited progress in NH₃ mitigation efforts to date, and the slow nature of behavioural change, stepping up efforts to control NH₃ emission must be a key priority for future policy development.

Full acknowledgements, including from the European Commission, European Space Agency, US EPA, other national funding sources and individuals are listed in the electronic supplementary material. This paper is a contribution to the International Nitrogen Initiative (INI) and to the UNECE Task Force on Reactive Nitrogen (UNECE).

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